

Density current parametrization at LMD and CNRM.

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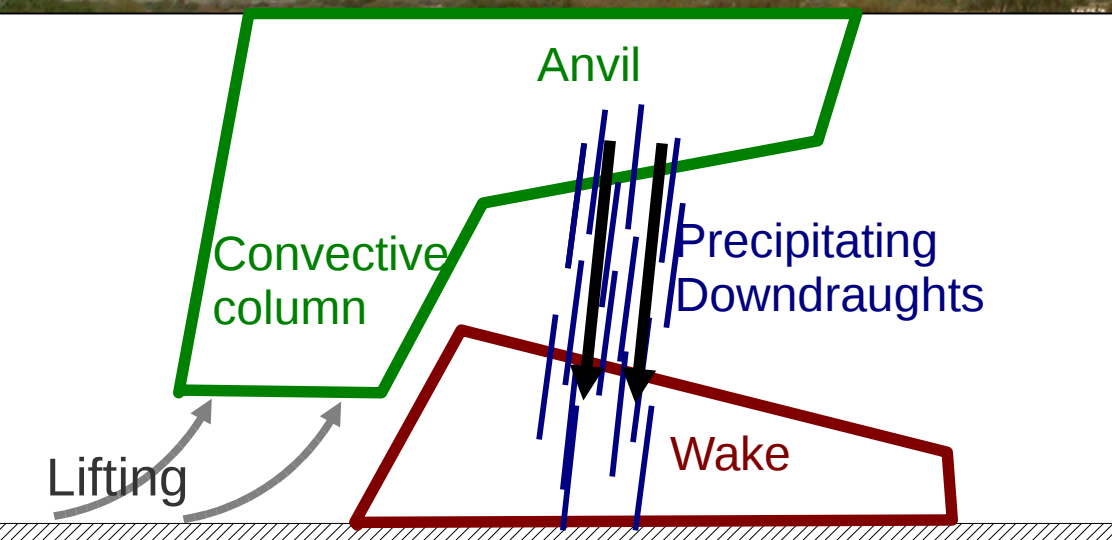
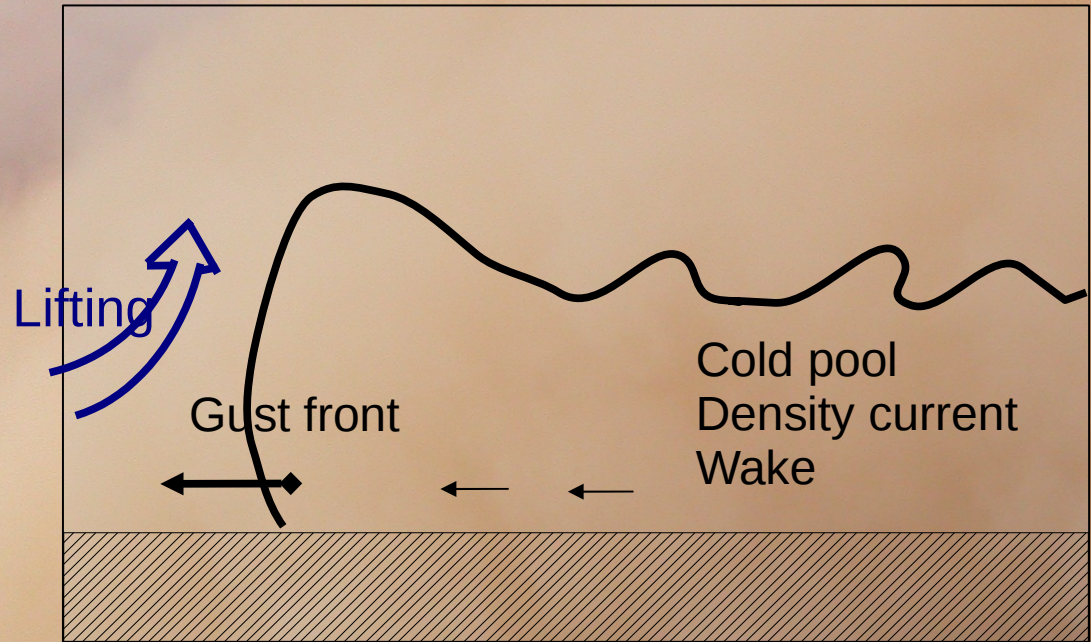
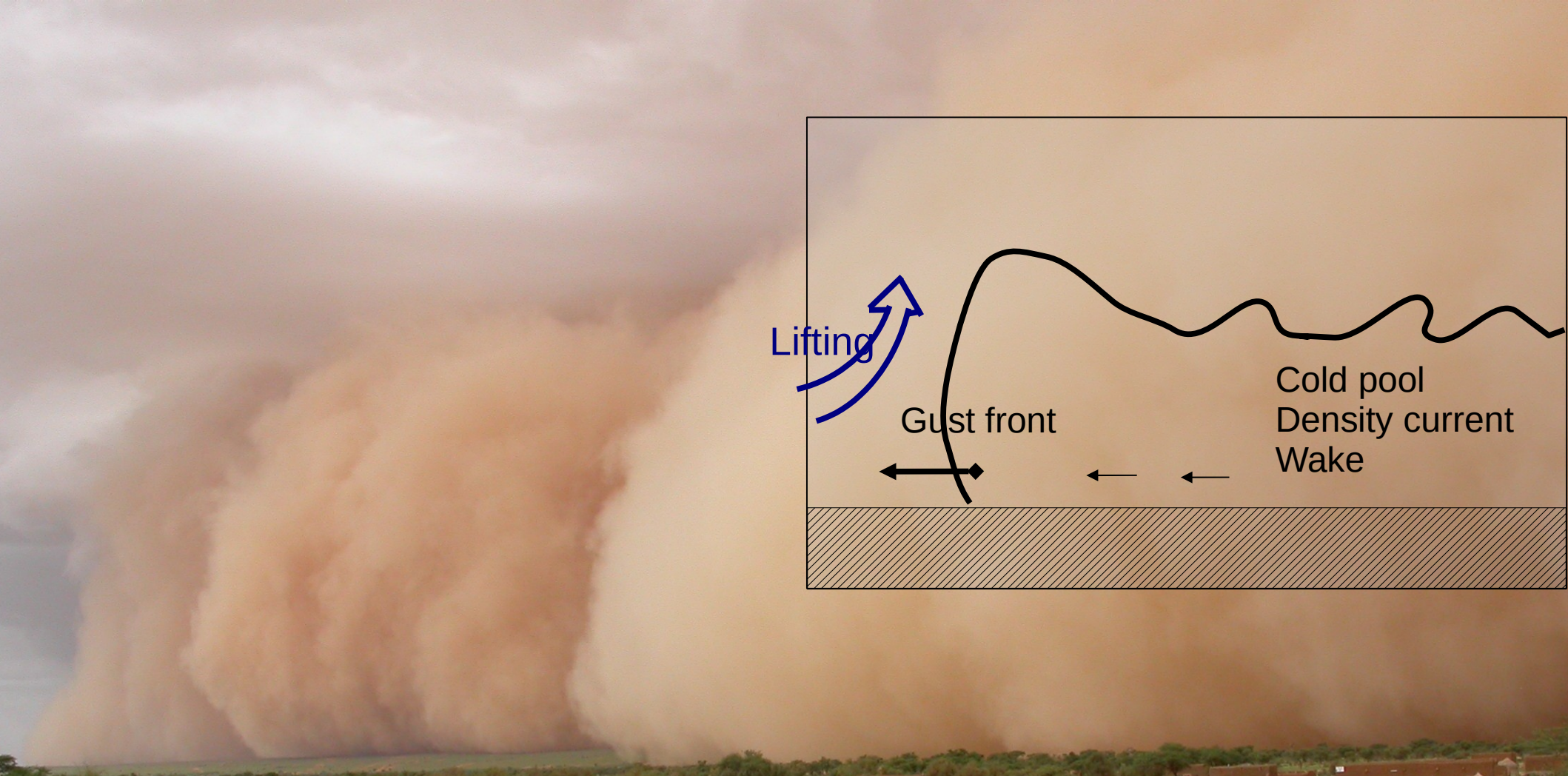
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Layout

- 1- Why parametrize density currents?
- 2- The LMD/CNRM density current (wake) parametrization.
- 3- Wakes in the LMDZ5 model.
- 4- LMDZ5 developments and present status.

Why parametrize density currents?

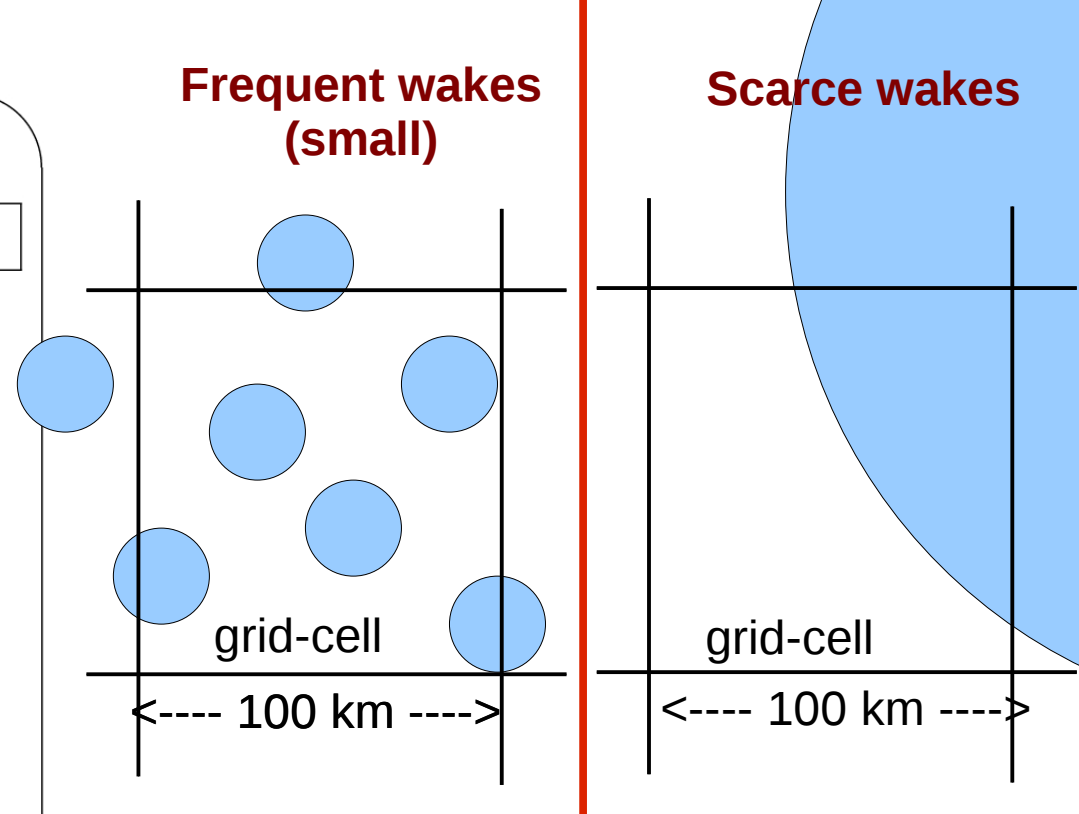
- In most GCMs deep convection is function only of Large Scale variables (quasi-equilibrium hypothesis) ==> elementary convective structures (Mesoscale Convective Systems, Squall Lines) do not exist in climate simulations.
- Observed self-sustaining behaviour of deep convection (diurnal cycle and propagation) shows that there is a memory somewhere.
- **Key hypothesis: wakes (= cold pools = density currents) are the key elements of this memory.**
- **LMDZ:** explicit representation of cold pools.



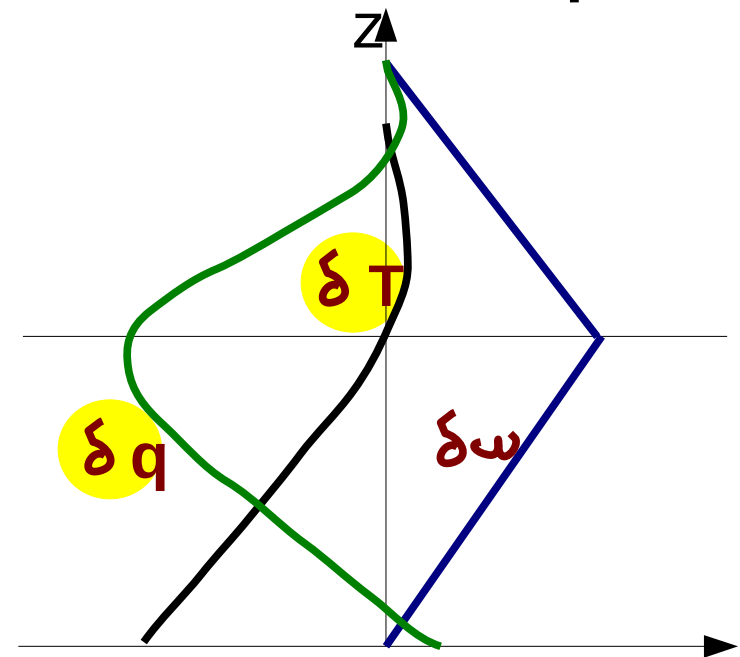
The density current (wake) parametrization

(Grandpeix and Lafore, JAS, 2010; Grandpeix et al., JAS 2010)

- Representation of a part of an infinite plane where identical cold pools (radius r , height h) are scattered with an homogeneous density D_{wk} .
- State variables : (i) surface fraction covered by the wakes $\sigma_w = \frac{S_w}{S_t}$ ($\sigma_w = \pi r^2 D_{wk}$), (ii) temperature and humidity differences (resp. $\delta\theta(p)$ and $\delta q(p)$) between wake and off-wake regions.
- Spreading speed : C_* such that $C_*^2 \simeq WAPE$ (WAKE Potential Energy); $WAKE = \int_{p_{top}}^{p_{surf}} R_d \delta T_v \frac{dp}{p}$
- Evolutions of $\delta\theta$ and δq profiles are given by conservation equations of mass, energy and water taking into account vertical advection, turbulence and phase changes.
- Turbulence and phase change terms are assumed to be given by the deep convection scheme.
- $\delta\omega$ profile is linear between the surface and the wake top (no mass exchange through the wake boundary); it goes back to 0 linearly between the wake top and an arbitrary altitude (about 4000 m).



Wake differential profiles



Wake variable tendencies

$$\left\{ \begin{array}{l}
 \partial_t \delta \theta = -\bar{\omega} \partial_p \delta \theta + \frac{\delta Q_1^{\text{cv}} + \delta Q_1^{\text{wk}}}{C_p} - \frac{k_{\text{gw}}}{\tau_{\text{gw}}} \delta \theta \\
 \text{where } \tau_{\text{gw}} = \frac{\sqrt{\sqrt{\sigma_w}(1-\sqrt{\sigma_w})}}{4Nz\sqrt{D_{\text{wk}}}} \\
 \text{is the damping time by gravity waves} \\
 \\
 \frac{\delta Q_1^{\text{wk}}}{C_p} = -\frac{e_w}{\sigma_w} \delta \theta \quad : \text{Entraînement} \\
 \\
 -\delta \omega \partial_p \bar{\theta} \quad : \text{differential advection of } \bar{\theta} \\
 \\
 -(1 - 2\sigma_w) \delta \omega \partial_p \delta \theta \quad : \text{differential advection of } \delta \theta
 \end{array} \right.$$

Water vapour equations are similar except for the gravity wave term :

$$(\partial_t \delta q_v)_{\text{gw}} = -\frac{k_{\text{gw}}}{\tau_{\text{gw}}} \delta \theta \frac{\partial_z [\bar{q}_v + (1 - 2\sigma_w) \delta q_v]}{\partial_z [\bar{\theta} + (1 - 2\sigma_w) \delta \theta]}$$

Wake variable tendencies

Differential heat source due to deep convection: precipitating downdraughts cool the wakes .

$$\partial_t \delta\theta = -\bar{\omega} \partial_p \delta\theta + \frac{\delta Q_1^{\text{cv}} + \delta Q_1^{\text{wk}}}{C_p} - \frac{k_{\text{gw}}}{\tau_{\text{gw}}} \delta\theta$$

$$\text{where } \tau_{\text{gw}} = \frac{\sqrt{\sqrt{\sigma_w}(1-\sqrt{\sigma_w})}}{4Nz\sqrt{D_{\text{wk}}}}$$

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$$\frac{\delta Q_1^{\text{wk}}}{C_p} = -\frac{e_w}{\sigma_w} \delta\theta \quad : \text{Entraînement}$$

$$-\delta\omega \partial_p \bar{\theta} \quad : \text{differential advection of } \bar{\theta}$$

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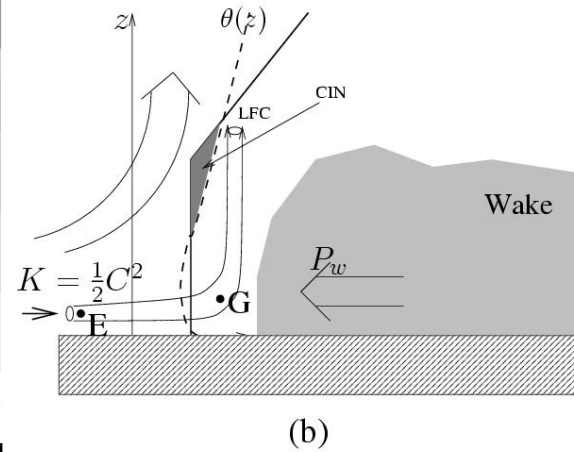
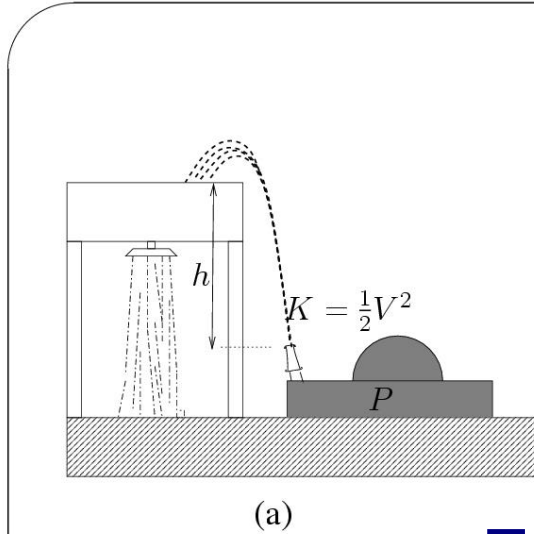
$$(\partial_t \delta q_v)_{\text{gw}} = -\frac{k_{\text{gw}}}{\tau_{\text{gw}}} \delta\theta \frac{\partial_z [\bar{q}_v + (1 - 2\sigma_w) \delta q_v]}{\partial_z [\bar{\theta} + (1 - 2\sigma_w) \delta\theta]}$$

Large scale variable tendencies

$$\left\{ \begin{array}{l} \partial_t \bar{\theta} = (\partial_t \bar{\theta})_{LS} + \frac{Q_R + Q_1^{\text{bl}} + Q_1^{\text{cv}} + Q_1^{\text{wk}}}{C_p} \\ \\ \frac{Q_1^{\text{wk}}}{C_p} = +(\partial_t \sigma_w - e_w) \delta\theta \quad : \text{Spreading \& entrainment (cooling)} \\ \\ \quad \quad \quad -\sigma_w(1 - \sigma_w) \delta\omega \partial_p \delta\theta \quad : \text{Differential advection (heating)} \end{array} \right.$$

with similar equations for water vapour.

Coupling convection with sub-cloud processes: ALE & ALP; 1



At least two variables:
 • the **Available Lifting Energy (ALE)**
 • the **Available Lifting Power (ALP)**.

Trigger

The **shower is triggered** when $K > gh$ ($K = \text{ALE} = \text{Available Lifting Energy}$).

Convection is triggered when the maximum kinetic energy K ($K = \text{ALE}$) of air impinging on the gust front exceeds the convective inhibition: **ALE > |CIN|**

Closure

The pump (power P) yields a mass flow rate M . A fraction k (the engine efficiency) of P is used by the stream.

The wakes provide a power P_w . A fraction k (the wake lifting efficiency) of P_w is used to lift draughts with mass flow rate M :

- overcoming inhibition \Rightarrow power $M |CIN|$
- velocity at LFC $= w_B \Rightarrow$ power $\frac{1}{2} M w_B^2$
- dissipation \Rightarrow power $\frac{3}{2} M w_B^2$

Closure: stream power $M K = k P$ (= ALP)

Closure: **$M (|CIN| + 2 w_B^2) = k P_w$** (= ALP)

Coupling convection with sub-cloud processes: ALE & ALP; 2

More generally, the convection parametrization is coupled to sub-cloud processes through the two variables: ALE (for the trigger) and ALP (for the closure).

$$\text{ALE} = \sup(\text{ALE}_{\text{PBL}}, \text{ALE}_{\text{ORO}}, \text{ALE}_{\text{WK}})$$

$$\text{ALP} = \text{ALP}_{\text{PBL}} + \text{ALP}_{\text{ORO}} + \text{ALP}_{\text{WK}}$$

ALE (Available Lifting Energy) (J/kg)

ALE = order of magnitude of the kinetic energy of the strongest updraughts (scale $\simeq km$).

- Boundary layer: $ALE \simeq (\frac{1}{2}w^2)_{max}, \simeq (\frac{1}{2}w^2)_{Thermals}$.
- Orography thermal effect: ALE estimated from the potential energy of the surface layer.
- Density currents: $ALE = \frac{1}{2}C^{*2}$, (C^* = gust front velocity).

ALP (Available Lifting Power) (W/m^2)

- PBL:

$$ALP = \frac{1}{2}\overline{\rho w^3} \quad (\simeq \text{qq } 0.01 \text{ W/m}^2)$$

- Density currents:

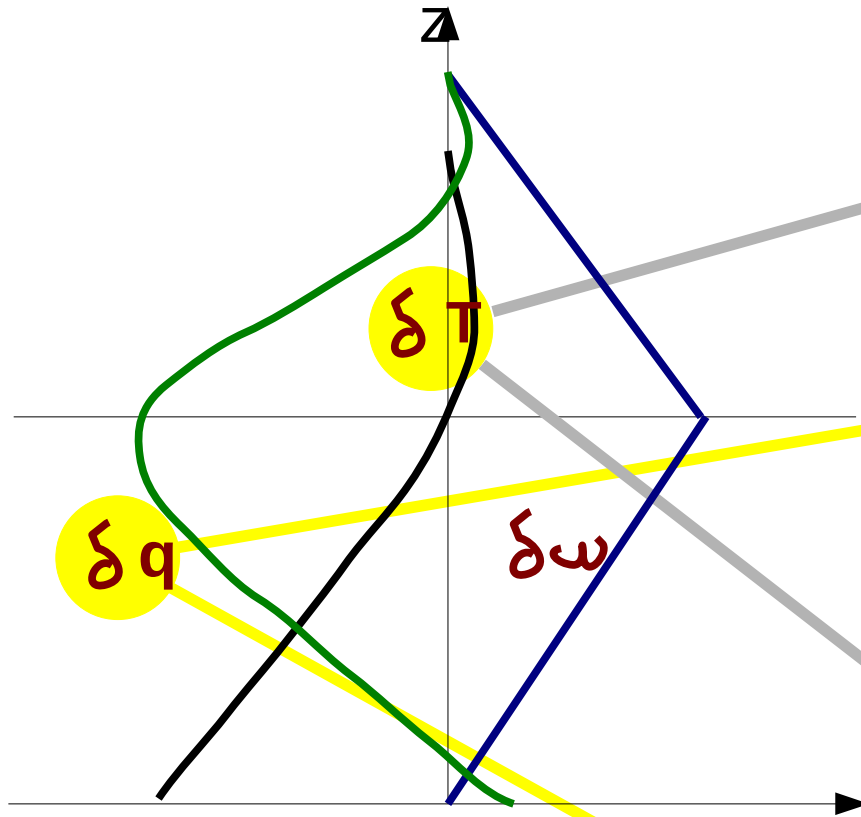
$$ALP = h_w \Gamma_w \frac{1}{2} \rho c^{*3} \quad (\Gamma_w = \text{gust frt lgth / unit area})$$

$(\simeq \text{qq } 0.1 \text{ W/m}^2)$

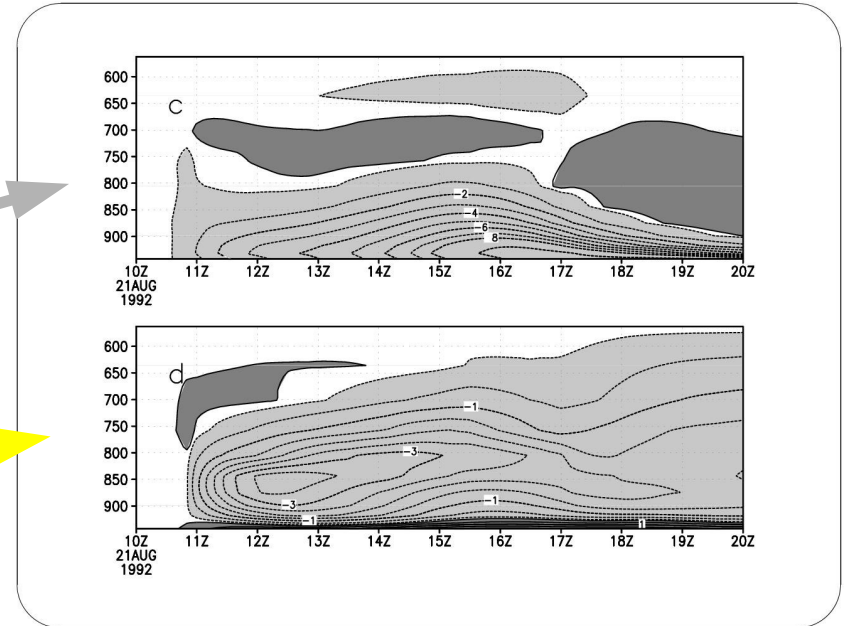
- Orography:

$$ALP = - \int_{top}^{base} \vec{D} \cdot \vec{V} dp \quad (\simeq \text{qq } 0.1 \text{ W/m}^2)$$

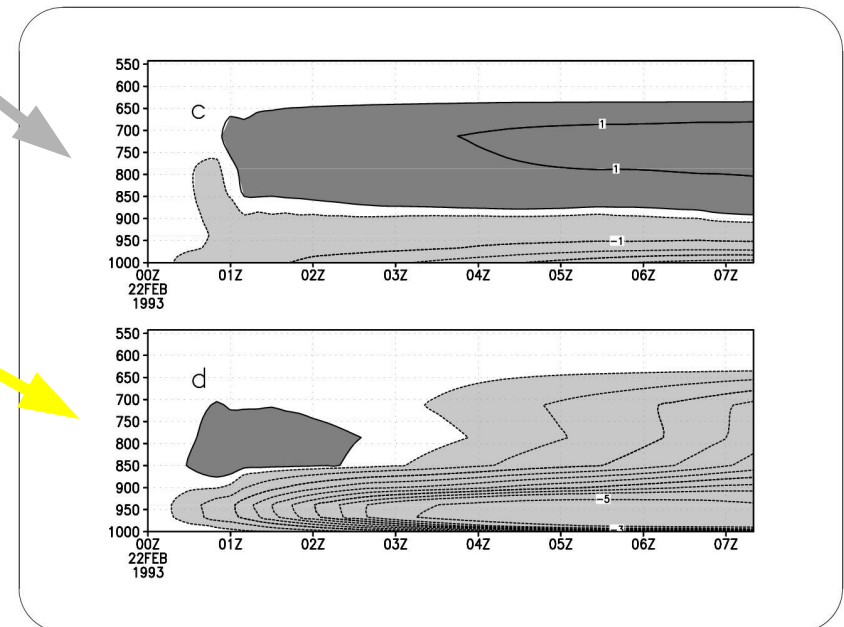
Simulated wake properties



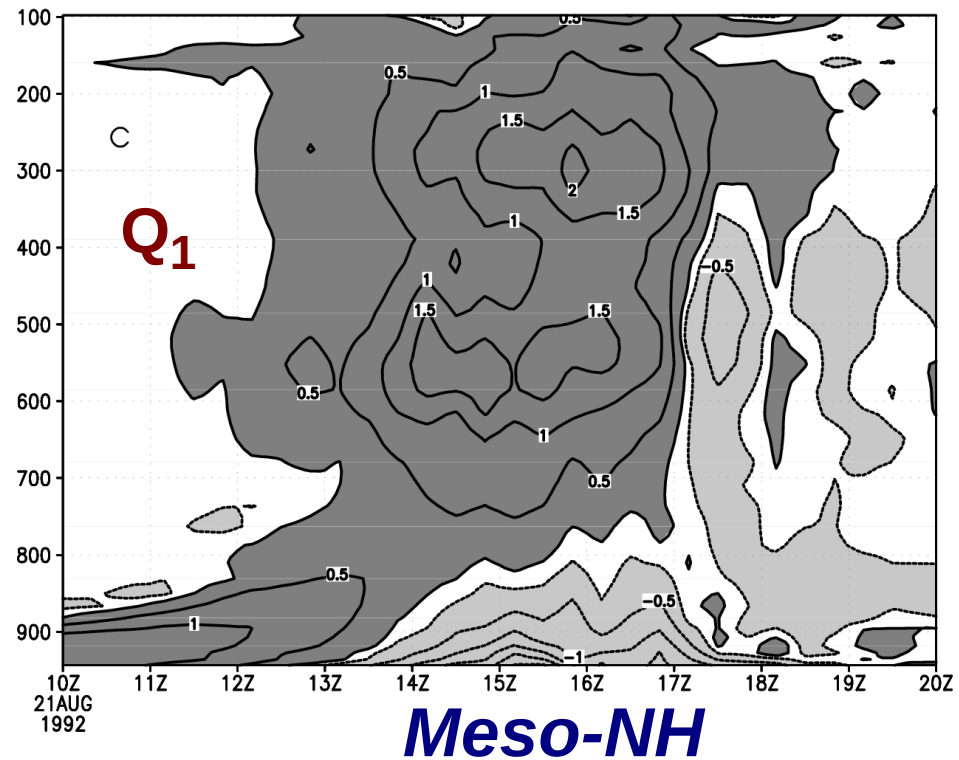
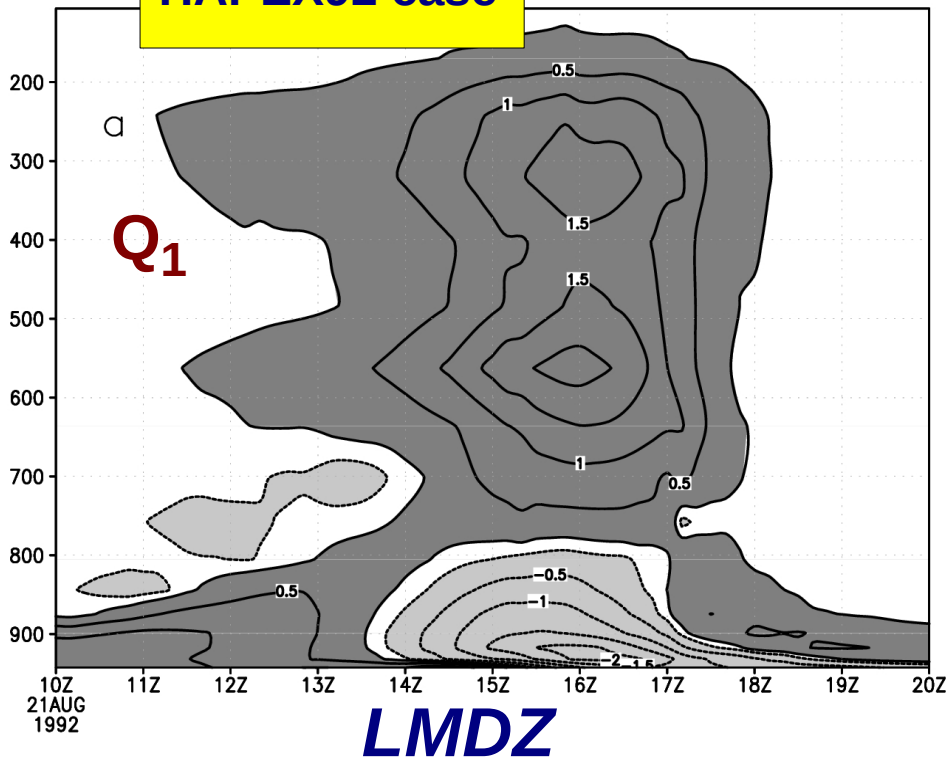
HAPEX92: 21 Aug 1992 squall line case



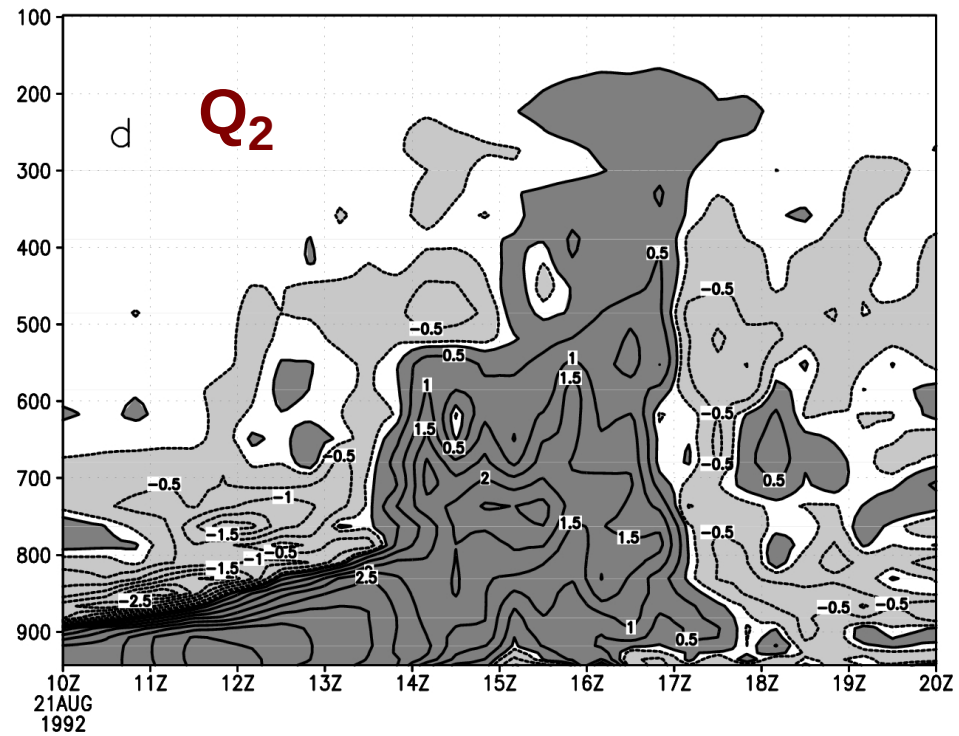
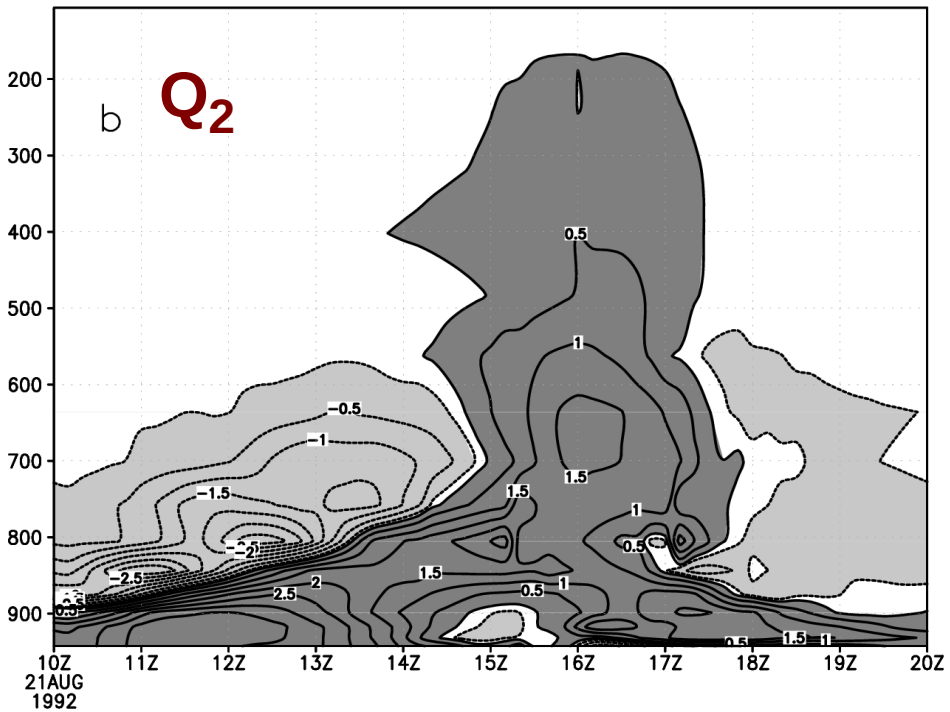
TOGA-COARE: 22 Feb 1993 squall line case



HAPEX92 case



Q2 (K/h) avec flux turbulents. Hapex (Initial)



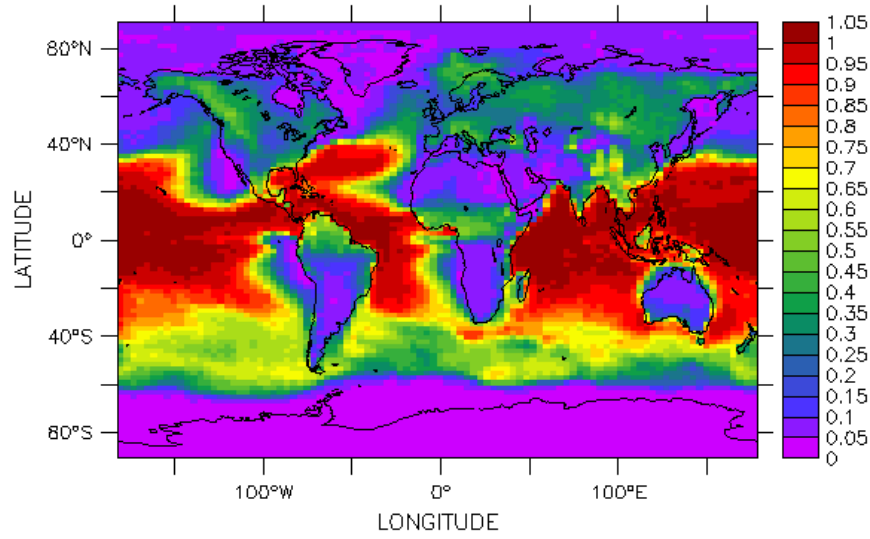
The wake parametrization present status

The wake parametrization is now routinely used in LMDZ5-NP (new physics) simulations.

In IPSLCM5B **CMIP5 simulations**, the Emanuel convective scheme is coupled with the [wake parametrization](#) and with the PBL scheme (cloudy thermals; Rio and Hourdin, 2008) via ALE and ALP.

T : 7

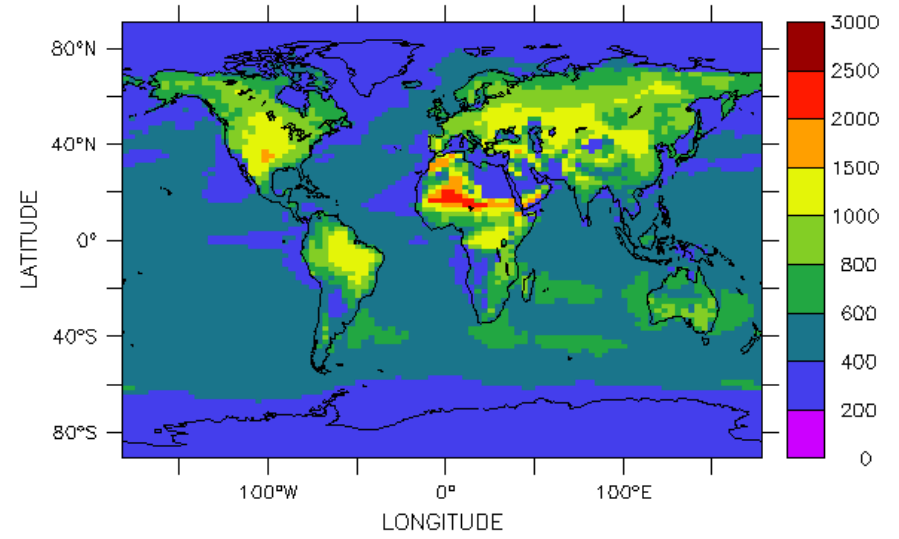
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Wake presence time, Jul

T : 7

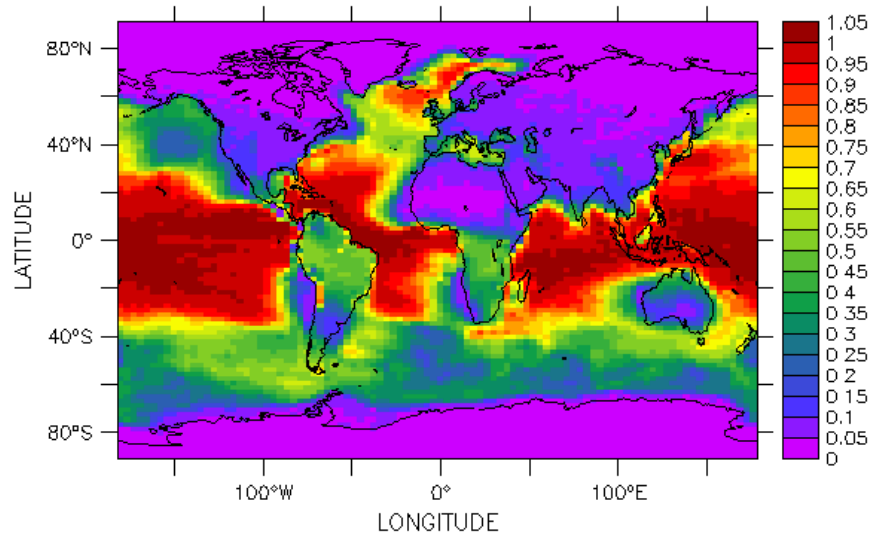
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Wake height, Jul

T : 2

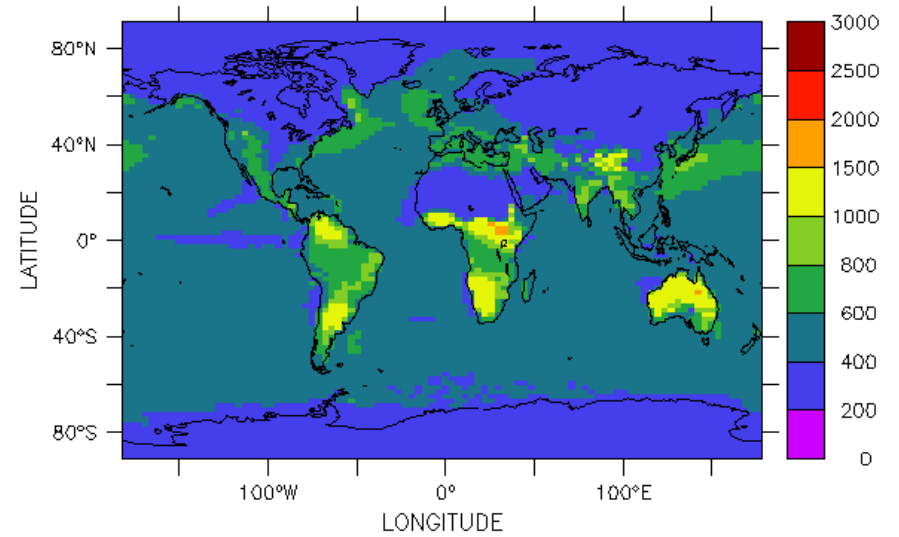
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Wake presence time, Feb

T : 2

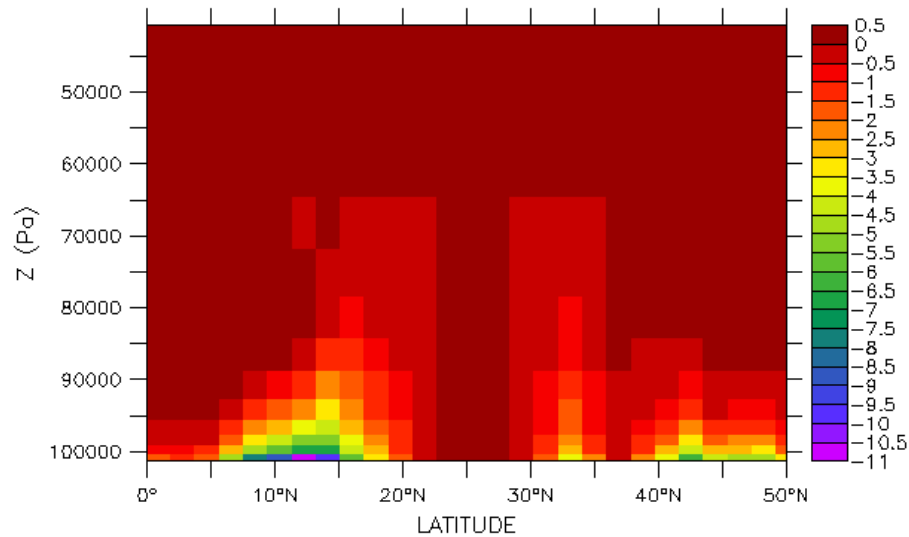
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Wake height, Feb

LONGITUDE : 0E
T : 8

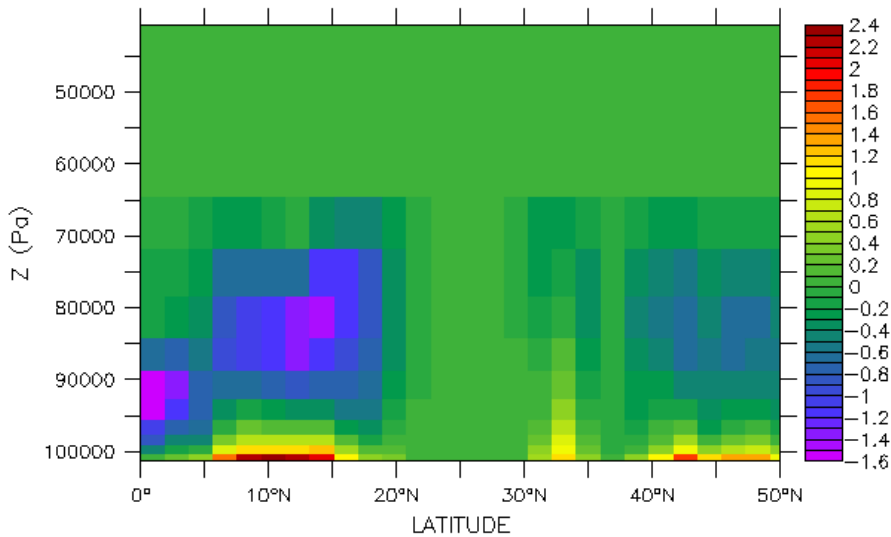
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WAKE_DELTAT*FW (K)

LONGITUDE : 0E
T : 8

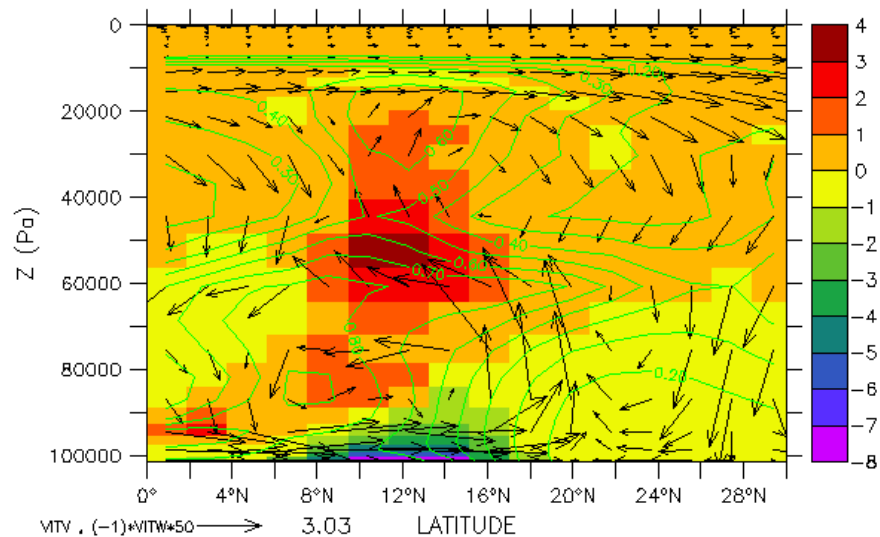
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WAKE_DELTAQ*FW (g/kg)

LONGITUDE : 0E
T : 8

DATA SET: NPv3.1_SE_1984_1991_1M_histmth



(DTCON+DTWAK)*86400

MTV, (-1)*MTW*50
CONTOUR: RHUM
CONTOUR: RHUM

And Now!

Ongoing developments:

- Improvement of some specific points of LMDZ parametrizations (splitting the PBL between the wake and the off-wake regions; introducing the ice thermodynamics).
- New parametrization of deep convection triggering (Nicolas Rochetin thesis).
- Introducing a representation of the wake population dynamic.

Need for satellite observations for the validation of deep convection triggering and for global analysis of cold pools.

Near future developments:

- Convection propagation.
- Anvil representation.

In the last two instances there are both technical issues (how do we do it?) and scientific issues (what are the processes to parametrize? **How to use satellite informations?**).